

On the use of spectral kernels

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Spectral kernel application

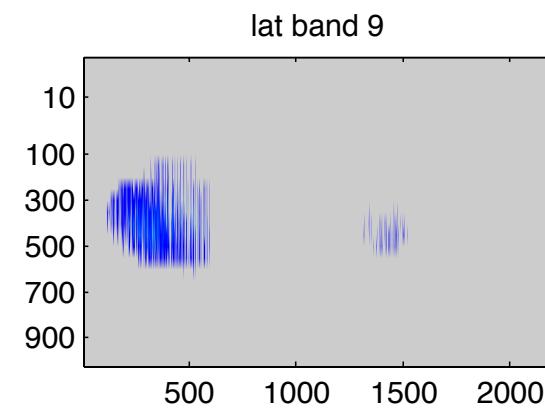
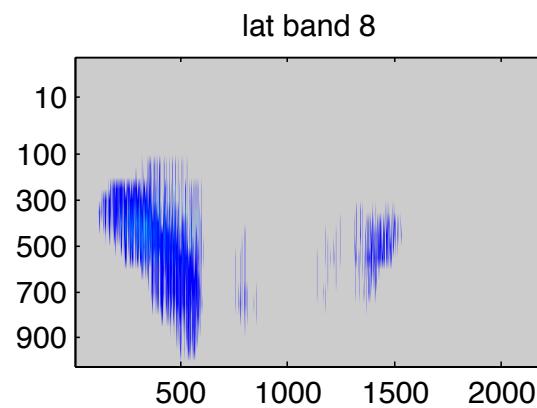
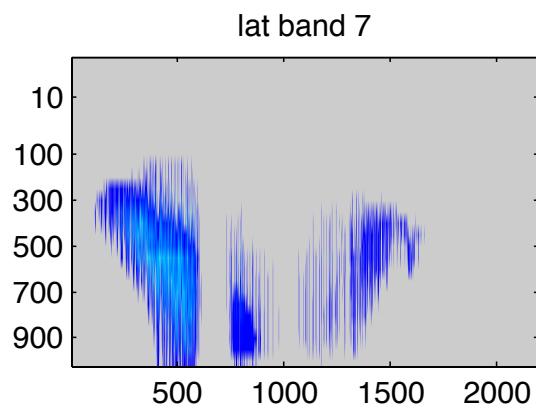
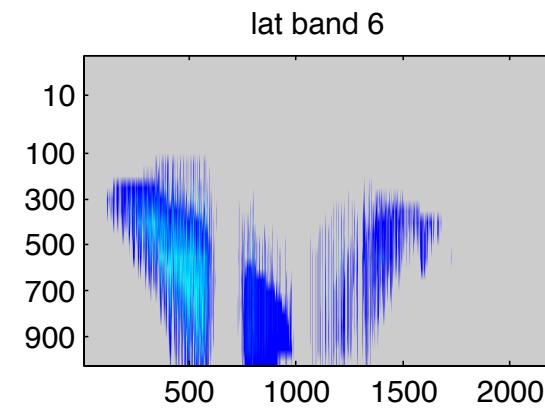
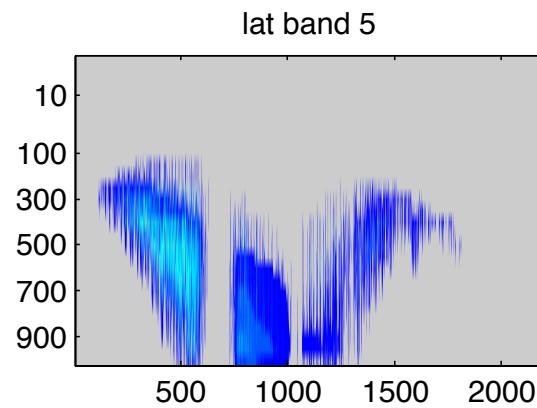
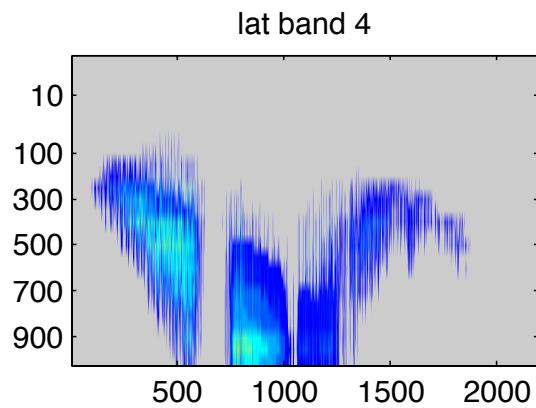
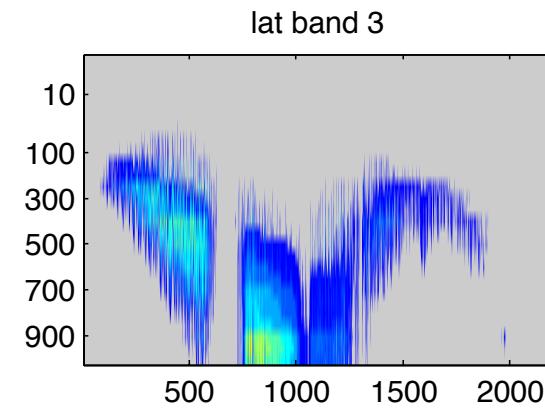
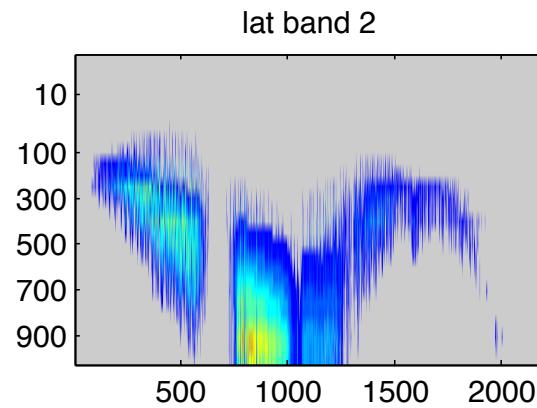
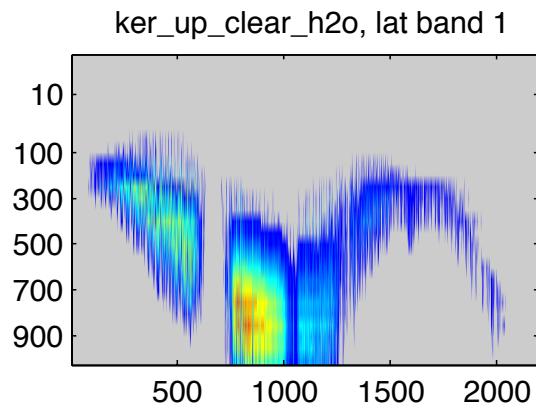
- Kernel (scaling) method
 - Spectral kernels
 - Radiative effect of water vapor
- Implication for radiative forcing in general
 - Curve of growth (spectroscopy) explanation vs. Radiative transfer explanation

Reference

Bani Shahabadi, M., and Y. Huang (2014), Logarithmic radiative effect of water vapor and spectral kernels, *J. Geophys. Res. Atmos.*, 119, 6000–6008, doi:10.1002/2014JD021623.

Zonal mean (equa. to pole) spectral kernels

Water vapor
clear-sky



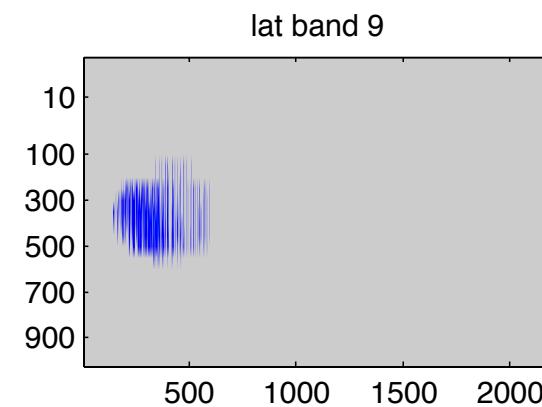
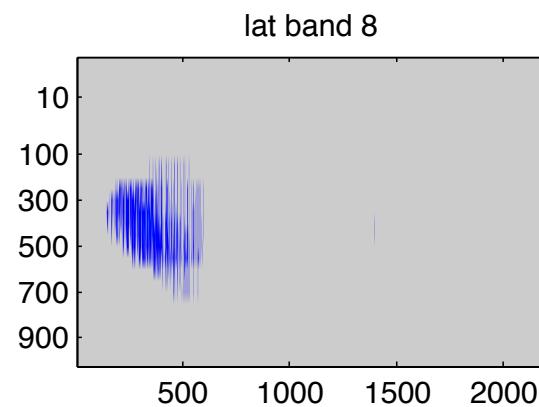
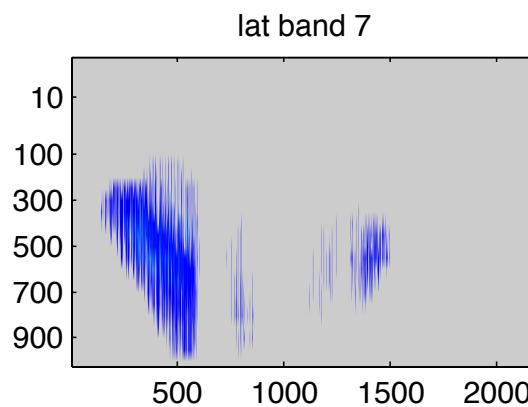
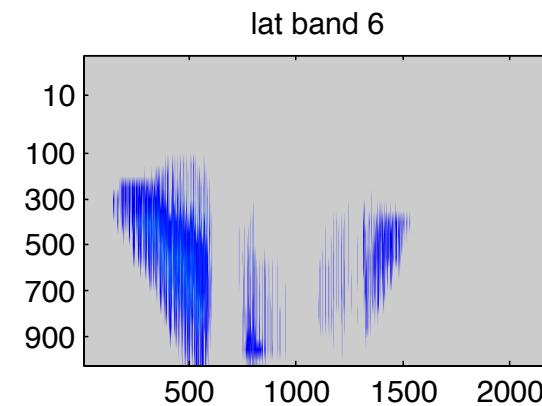
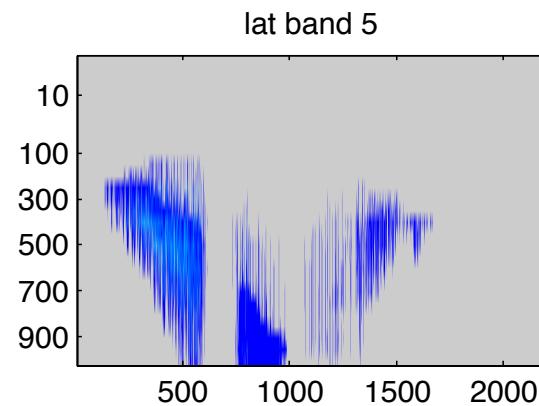
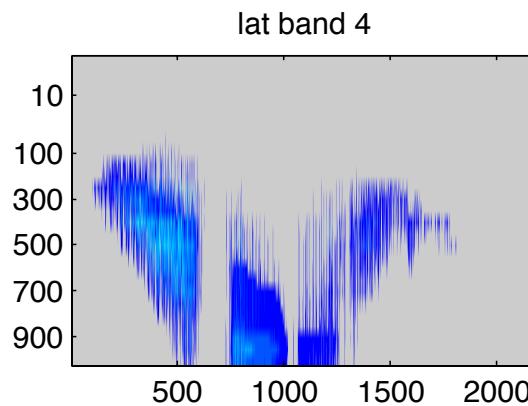
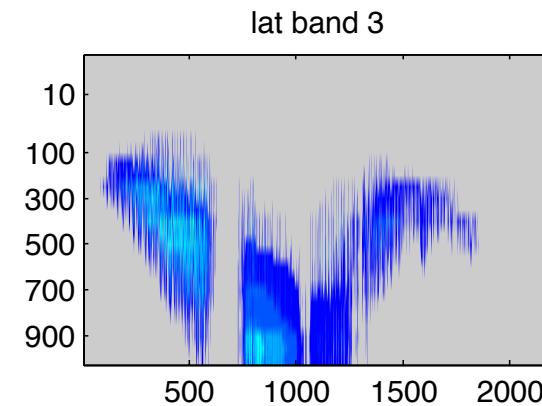
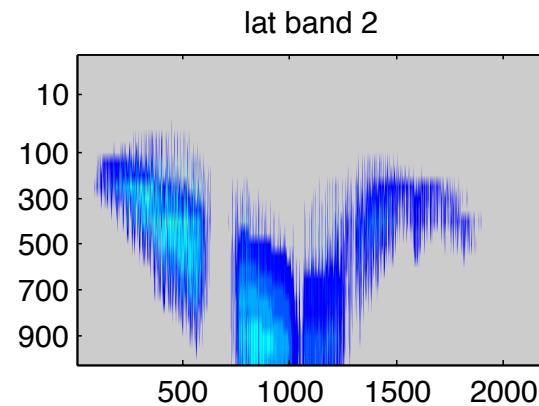
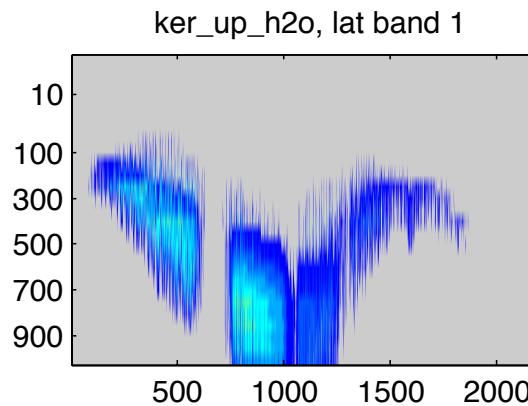
W m^{-2}
 $/ \text{cm}^{-1}$
 $/ \Delta\chi$

$\times 10^{-4}$

A vertical color bar indicating the magnitude of the spectral kernel values. The scale ranges from 0 (dark blue) to 8 (dark red), with intermediate ticks at 2, 4, and 6. The values are multiplied by 10^{-4} , as indicated by the label.

Zonal mean (equa. to pole) spectral kernels

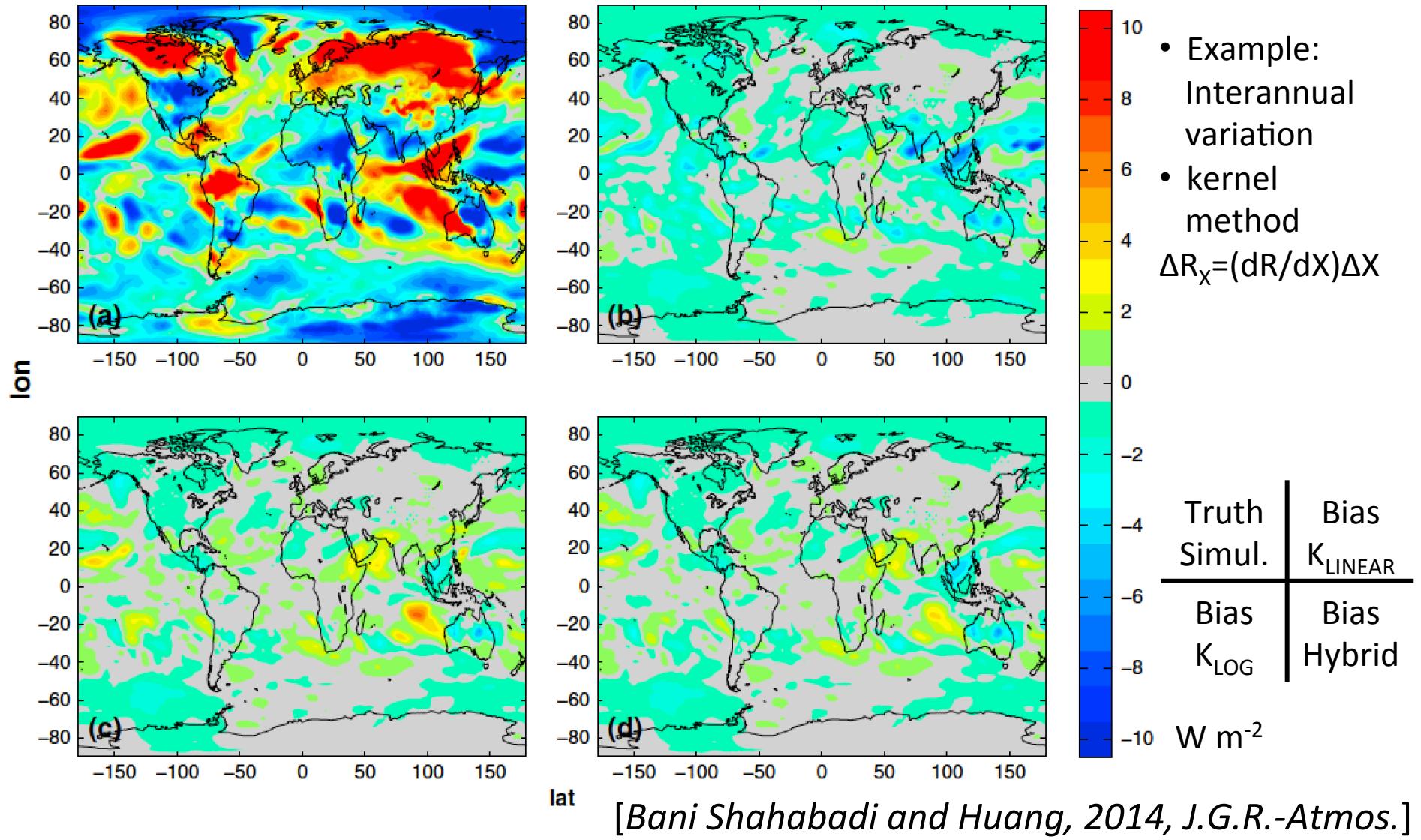
Water vapor
all-sky



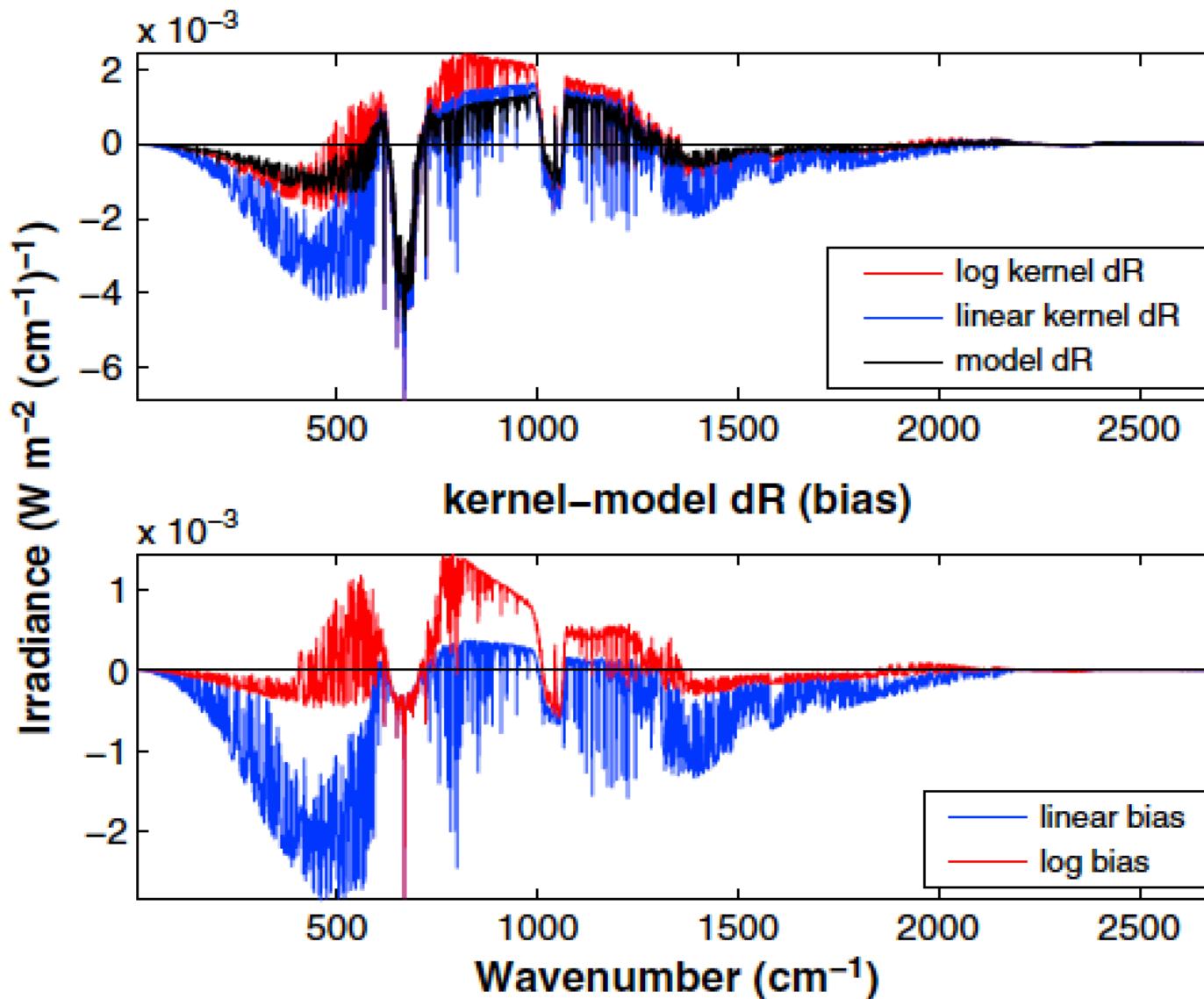
W m^{-2}
 $/ \text{cm}^{-1}$
 $/ \Delta \chi$

$\times 10^{-4}$
0 2 4 6 8

Model-simulated vs. Kernel-analyzed dOLR



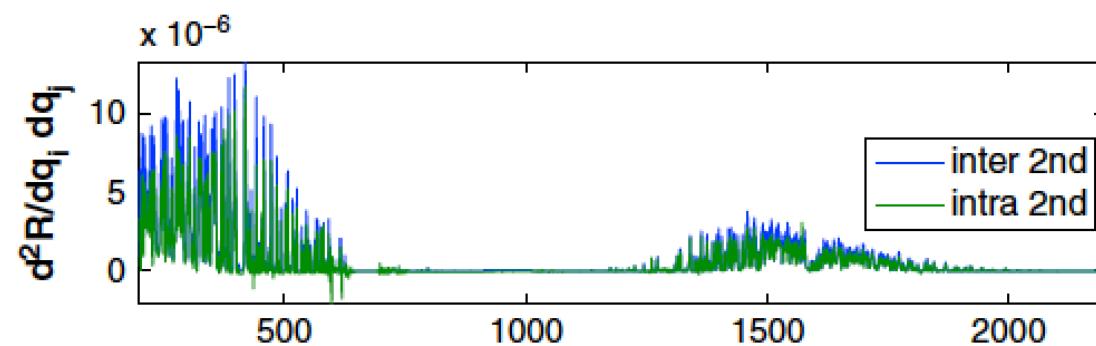
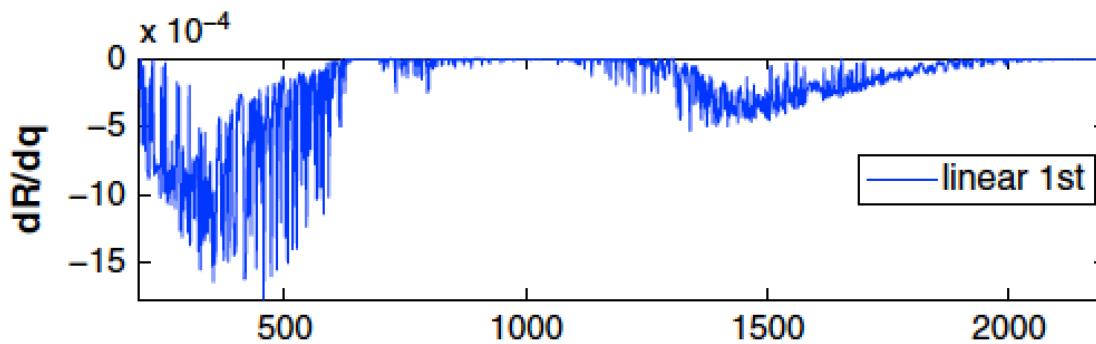
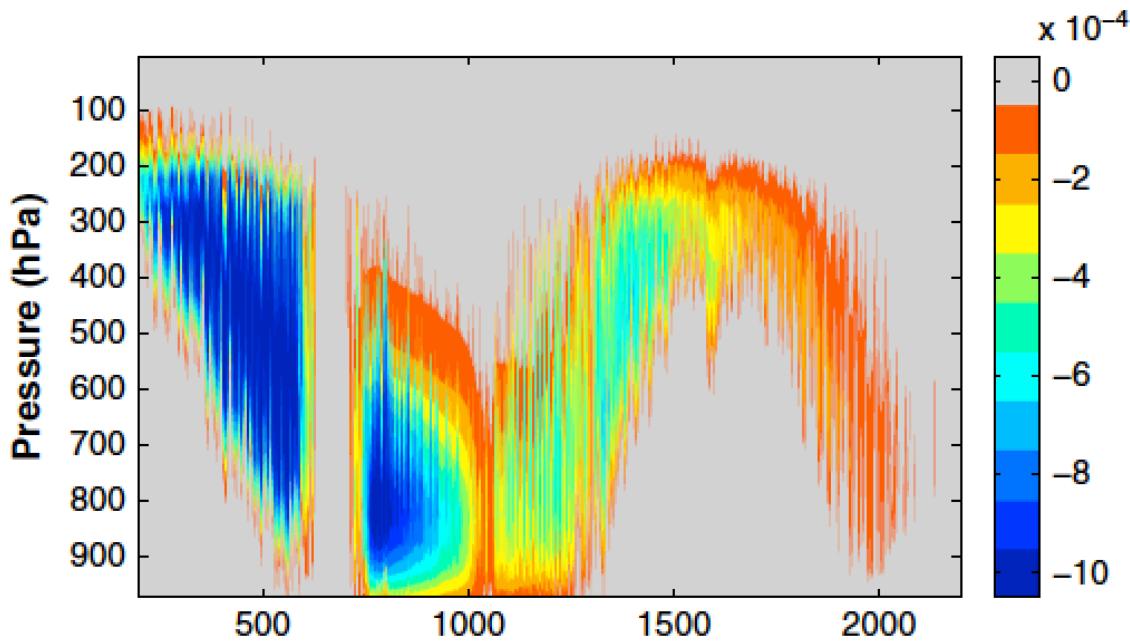
Model-simulated vs. Kernel-analyzed spectral dOLR



- Interannual tropical mean change
 - kernel method
- $$\Delta R_X = (dR/dX)\Delta X$$

Concerning H₂O:
Linear scaling:
 $\Delta X = \Delta(q)$

Log-scaling:
 $\Delta X = \Delta(\log(q))$

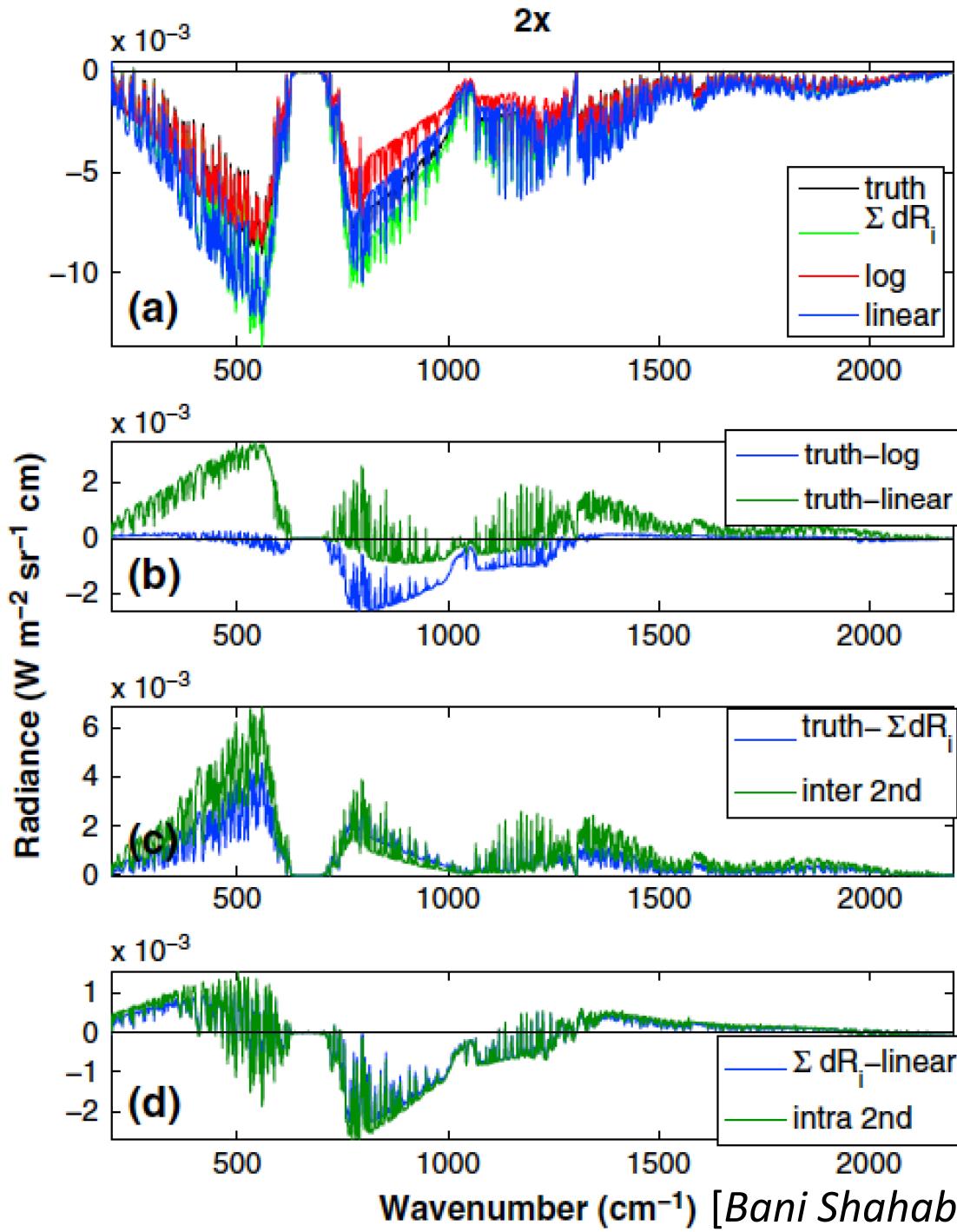


Wavenumber (cm^{-1}) [Bani Shahabadi and Huang, 2014, J.G.R.-Atmos.]

1st vs. 2nd-order kernel

- Log behaviour, i.e., reduced sensitivity, can be understood as inter-layer and intra-layer effects.

$$\Delta R = \sum_i \frac{\partial R}{\partial q^i} \Delta q^i + \frac{1}{2} \sum_i \frac{\partial^2 R}{\partial q^i \partial q^i} (\Delta q^i)^2 + \sum_{i \neq j} \frac{\partial^2 R}{\partial q^i \partial q^j} \Delta q^i \Delta q^j$$



1st vs. 2nd-order kernel

- Case: uniform 2xH₂O
- Inter-layer > intra-layer for H₂O band
- Inter-layer offsets intra-layer

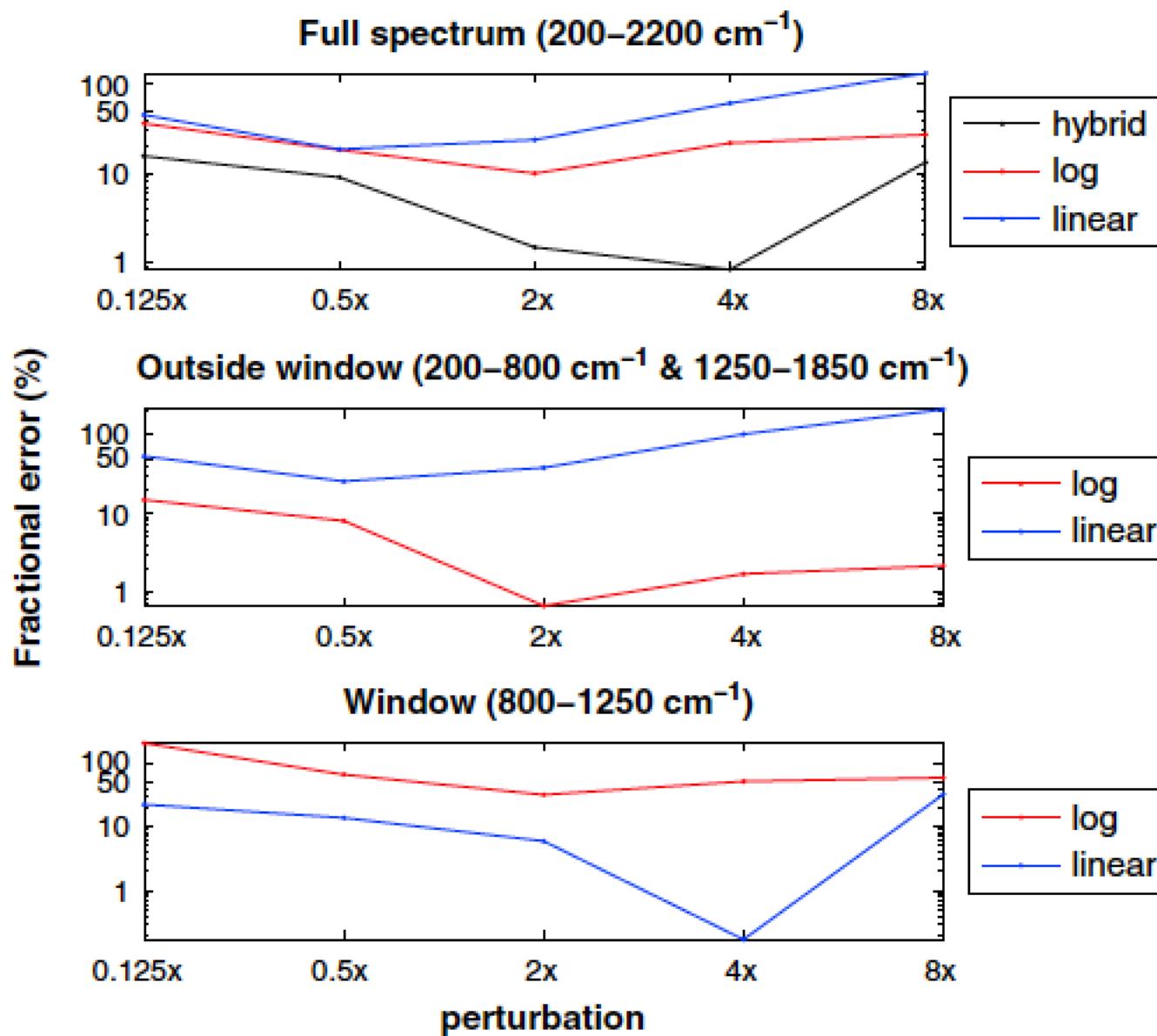
$$\Delta R = \sum_i \frac{\partial R}{\partial q^i} \Delta q^i$$

$$+ \frac{1}{2} \sum_i \frac{\partial^2 R}{\partial q^i \partial q^i} (\Delta q^i)^2$$

$$+ \sum_{i \neq j} \frac{\partial^2 R}{\partial q^i \partial q^j} \Delta q^i \Delta q^j$$

[Bani Shahabadi and Huang, 2014, J.G.R.-Atmos.]

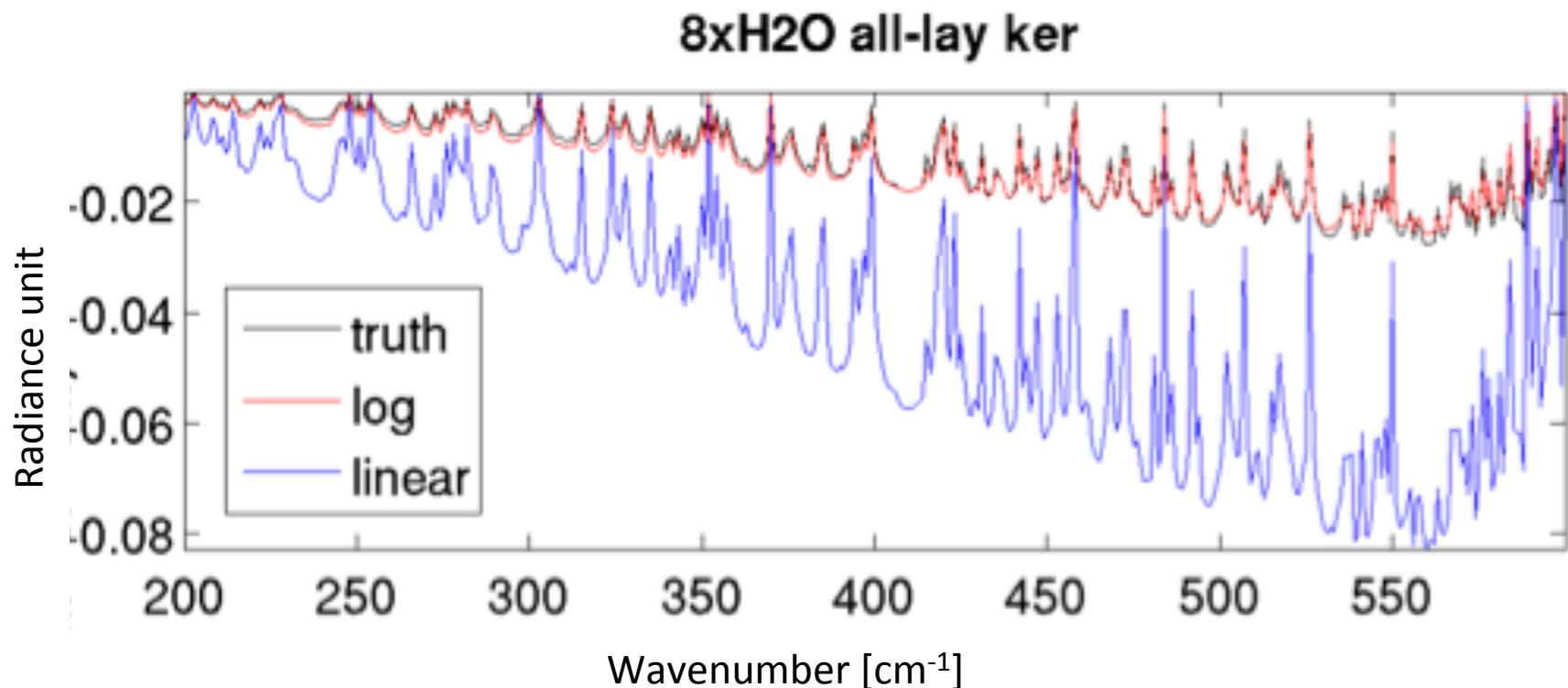
Hybrid scaling



Best performance can be obtained through:

- Absorption band: log-scaling
- Window: linear-scaling

Logarithmic relationship holds
even for monochromatic radiance!



Monochromatic!

Truth:

$$\Delta R(8x) = R(8x) - R(1x)$$

Calculated using LBL RT model

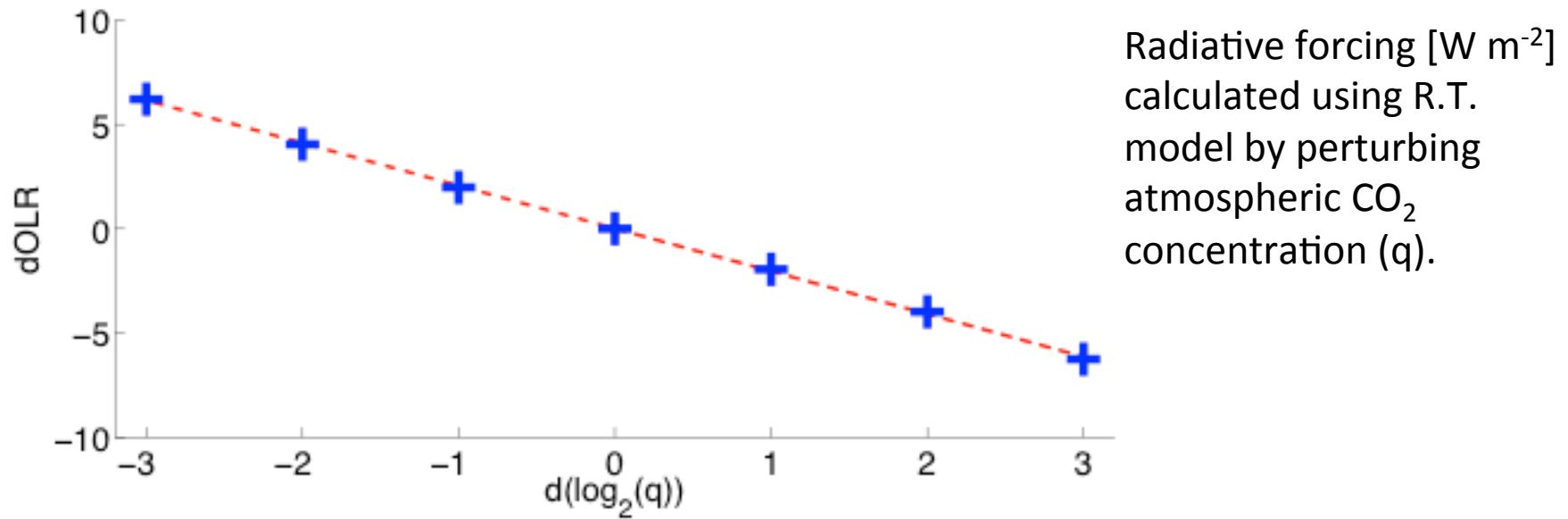
Log scaling:

$$\Delta R(8x) = \Delta R(1.1x) * \log(8)/\log(1.1)$$

Linear scaling:

$$\Delta R(8x) = \Delta R(1.1x) * (8-1)/0.1$$

Logarithmic relationship

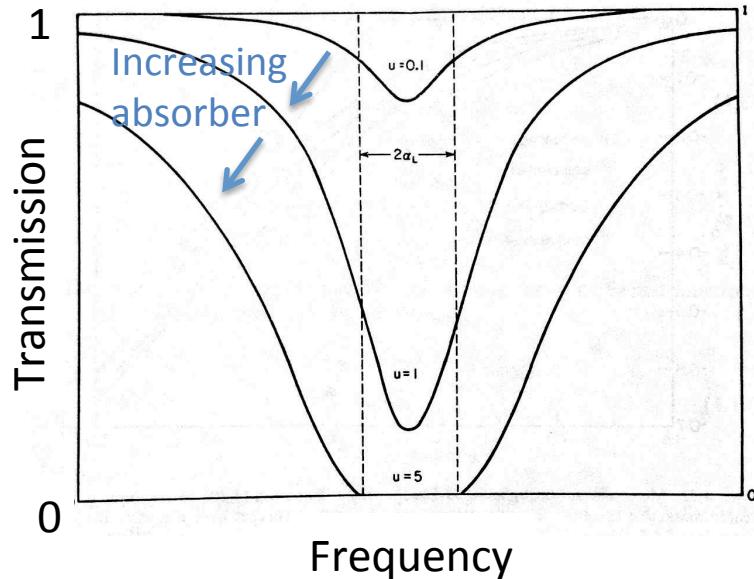


- Log relationship can be verified with any RT model
 - Log estimation formula widely adopted
 - $F = F_0 \log(q/q_0)$
- [IPCC AR1,2,3,...; Wikipedia, ...]

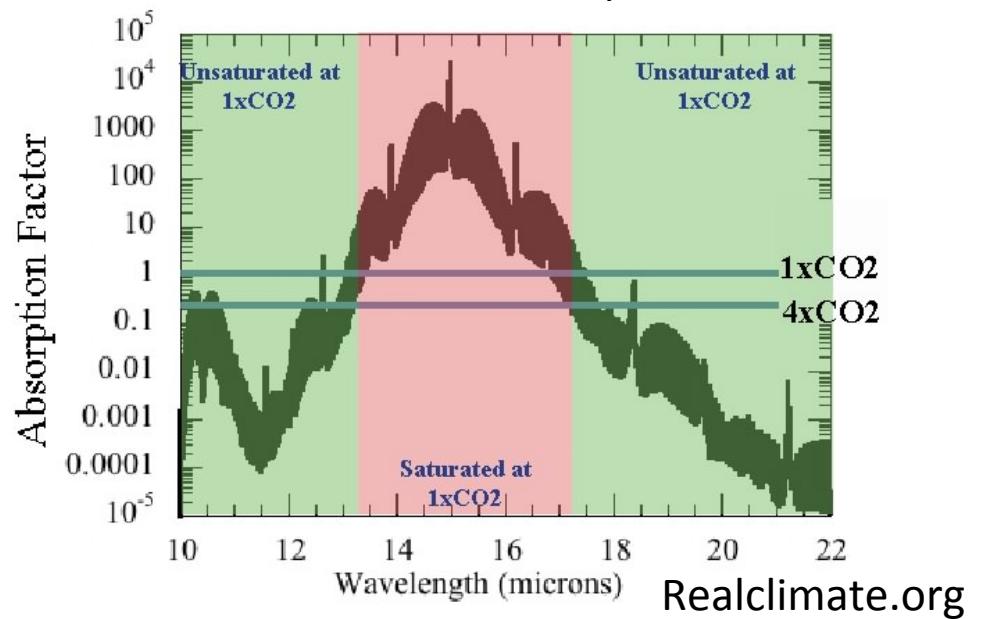
Cause of logarithmic relationship

Answer 1: it is due to spectroscopy

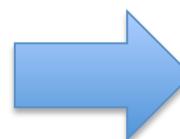
Saturation of an absorption line



Saturation of an absorption band

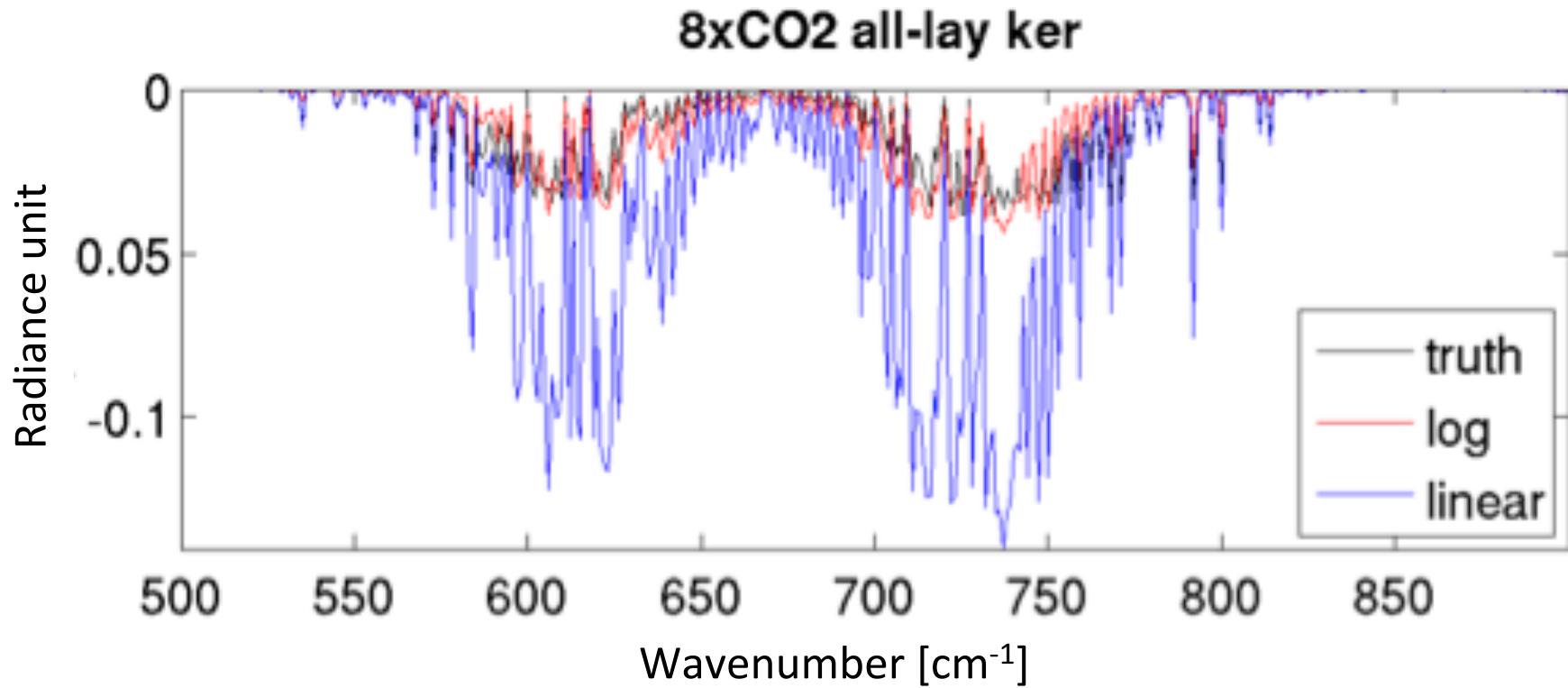


Radiative forcing \propto increased absorption
 \propto saturation from line center to wing: “curve of growth” theorem [Goody&Yung 1989];
from band center to wing [Pierrehumbert 2010]



The log relationship applies to spectrally integrated (broadband) radiation flux.

Counterevidence: Logarithmic relationship holds even for monochromatic radiance!



CO₂ case.

Truth:

$$\Delta R(8x) = R(8x) - R(1x)$$

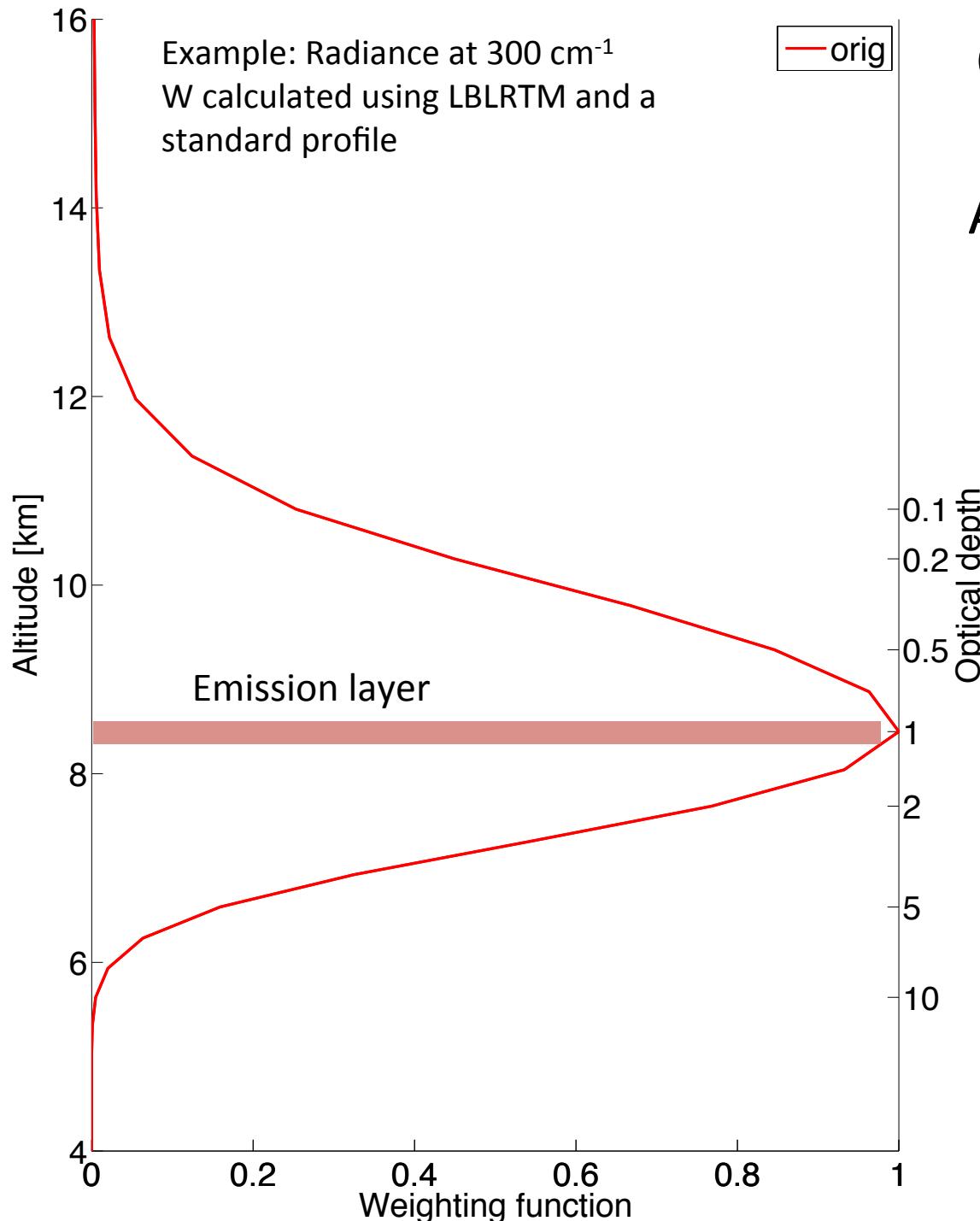
Calculated using LBL RT model

Log scaling:

$$\Delta R(8x) = \Delta R(1.1x) * \log(8)/\log(1.1)$$

Linear scaling:

$$\Delta R(8x) = \Delta R(1.1x) * (8-1)/0.1$$



Cause of logarithmic relationship

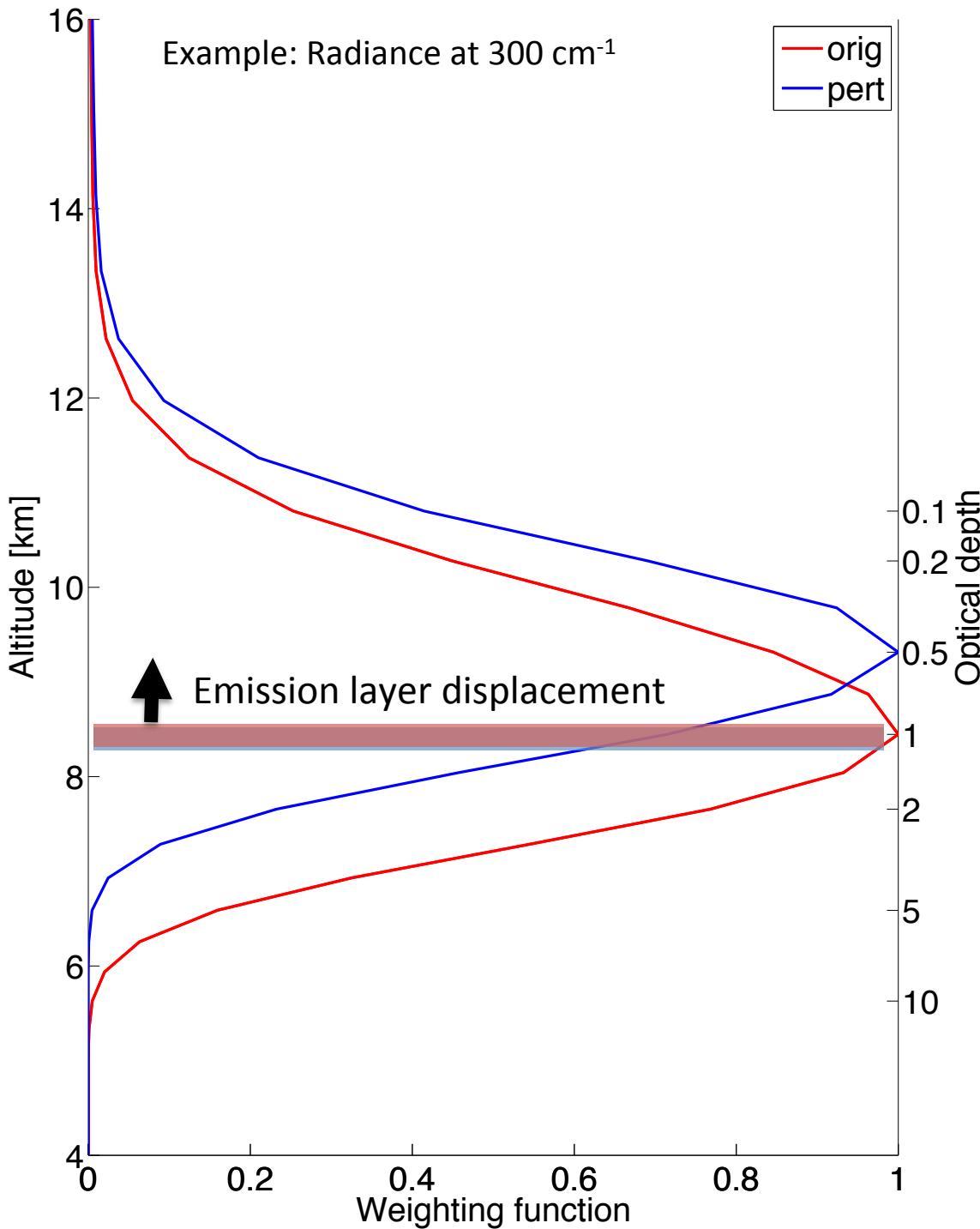
Answer 2: it is due to radiative transfer

- Emission layer displacement model

Solution to non-scattering R.T. Eq. can be generalized as:
 $R = \sum \{W_i * B_i\}$

W_i : weighting function for layer i , a function of optical depth τ measured from TOA to layer i .

B_i : Planck function of layer temperature (T_i)



Emission layer displacement model

Solution to non-scattering R.T. Eq. can be generalized as:

$$R = \sum \{W_i * B_i\}$$

Perturbation of absorber amount ($\alpha \times q$) equivalently displaces all the contributing layers to higher altitudes.

As $W = W(\tau)$ and $\tau \propto q$, each emission layer is displaced from τ to $\tau' = \tau/\alpha$.

Given $T = T_0 - z * \Gamma$, it can be shown $B \propto \log(\tau)$, and thus $B(\tau') - B(\tau) \propto \log(\alpha)$

[Huang and Bani Shahabadi, *J. Atm. Sci.*, under review]

Summary

- Spectral kernels are developed and ready for climate diagnoses
 - Log- and linear-scaling suit different spectral regions
 - Wise to apply hybrid scaling
- Why logarithmic in the absorption band?
 - Log-dependence:
 - Holds for monochromatic radiance (not only broadband flux)
 - Results from radiative transfer (not only spectroscopy)
“Emission layer displacement model”
 - Key conditions:
 - saturated absorption;
 - lapse rate;
 - coherent $dX \rightarrow$ fingerprint configuration